

Broadcasting Authority of Ireland Seminar

Climate Change: The Science

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Learning outcomes

By the end of this seminar, you should be able to:

- 1. Understand how climate is changing
- 2. Explain radiation and energy balance
- 3. Discuss the role of climate modelling in climate science
- 4. Describe the carbon cycle
- 5. Understand the impacts of climate forcings, feedbacks and climate sensitivity on climate change
- 6. Causes of climate change
- 7. Examine future climate change predictions, including for Ireland

1. Understand how climate is changing



Climate and weather

- Climate: average weather in a given area over a longer period of time. The classical period used for describing a climate is a 30 years (World Meteorological Organization).
- Weather: weather is short-term (minutes to days) variations in the atmosphere.
- Mark Twain: "Climate is what you expect; weather is what you get"

Global annual average temperature anomaly



Figure 2.2 Global annual average temperature anomaly (°C); the gray line is the annual average, and the blue line is a smoothed time series. Anomalies are calculated relative to the 1951–1980 average.

Data are from Berkeley Earth (http://berkeleyearth.org/data/, retrieved September 27, 2020).







(b) the warming as a function of latitude. Warming is calculated as the difference between the 1850-1900 average and

2009-2018 average.

Source: Dessler (2021)

Data are from Berkeley Earth (http://berkeleyearth.org/data/, retrieved October 14, 2020).

Distribution of modern warming

Satellite observations



Figure 2.4 Satellite measurements of the global monthly average temperature anomaly (°C) (blue line). The orange line is annual average temperature (°C) from the surface thermometer record plotted in <u>Figures 2.2</u> and <u>2.3</u>. Anomalies in this plot are relative to the 1981–2010 period.

Satellite data were obtained from the University of Alabama, Huntsville (<u>http://vortex.nsstc.uah.edu/data/msu/v6.0/tlt/uahncdc_lt_6.0.txt</u>, retrieved September 27, 2020).

Figure 2.6

Ice observations - Glaciers



(a) Global average annual change in mass of the world's glaciers, in tonnes/m²/year.



(b) Global average cumulative mass change, in tonnes/m². Given that ice has a density of about 0.9 tonnes/m³, a loss of 1 tonne of ice/m² is equal to the loss of about 1.1 m of glacier ice thickness.

Data from WGMS (2020) (obtained from https://wgms.ch/global-glacier-state/, retrieved November 26, 2020).



Figure 2.7 Arctic sea-ice area (in millions of square kilometers) in September of each year.

Data from Fetterer et al. (2017). Sea Ice Index, Version 3.0 [Data sets ID NSIDC-0081 and ID NSIDC-0051]. Boulder, Colorado USA: National Snow and Ice Data Center. doi: <u>https://doi.org/10.7265/N5K072F8</u> (accessed on November 26, 2020).

Ocean temperatures



Source: NOAA (2020)

Sea level



Figure 2.9 Change over time in global-average sea level, in millimeters. The seasonal cycle has been removed.

Data are described by Nerem et al. (2010) (obtained from http://sealevel.colorado.edu/, retrieved on November 26, 2020).

Climate change hiatus?



Figure 2.5 A plot of monthly and global average surface temperature (°C) from the surface thermometer record (blue line) along with short-term trend lines (red lines). This figure is an adaptation of SkepticalScience's escalator plot (<u>www.skepticalscience.com/graphics.php?g=47</u>).

Data are from Berkeley Earth (http://berkeleyearth.org/data/, retrieved September 28, 2020).

Earth's long-term climate record



Figure 2.13 Temperature anomaly of the southern polar region (blue line) over the past 410,000 years, relative to today's temperature, constructed from an Antarctic ice core. Carbon dioxide (orange line) is from air bubbles trapped in the ice

(data from Petit et al. (2000); downloaded from <u>https://cdiac.ess-dive.lbl.gov/climate/paleo/paleo_table.html</u>, accessed September 29, 2020).

Putting it all together

		IPCC limate change
CI Th	imate Change 202 e Physical Science Bas Summary for Policymakers	21 sis
wgi	Working Group I contribution to the Sixth Assessment Report of the Intergovernmental Panel on Climate Change	

IPCC assessment	Attribution of cause
First (1990)	The size of this warming is broadly consistent with prediction of climate models, but it is also of the same magnitude as natural climate variability. Thus the observed increase could be largely due to this natural variability
Second (1995)	The balance of evidence suggests a discernible human influence on the climate
Third (2001)	Most of the observed warming over the past 50 years is likely to have been due to the increase in greenhouse gas concentrations
Fourth (2007)	Most of the observed increase in globally averaged temperatures since the mid-twentieth century is very likely due to the observed increase in anthropogenic greenhouse gas concentrations
Fifth (2013)	It is extremely likely that more than half of the observed increase in global average surface temperature from 1951 to 2010 was caused by the anthropogenic increase in greenhouse gas concentrations and other anthropogenic forcings together
Sixth (2021)	Unequivocal that human influence has warmed the atmosphere, ocean and land. Widespread and rapid changes in the atmosphere, ocean, cryosphere and biosphere have occurred.
	Human-induced climate change is already affecting many weather and climate extremes in every region across the globe. Evidence of observed changes in extremes such as heatwaves, heavy precipitation, droughts, and tropical cyclones, and, in particular, their attribution to human influence, has strengthened since 5th assessment

2. Radiation and energy balance

Photons







Electromagnetic radiation and the atmosphere

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Energy Balance – First Law of Thermodynamics

Table 3.1 Energy balance summary

Condition	Effect on internal energy	Effect on temperature
If $E_{in} > E_{out}$ for an object	Internal energy of that object increases	Temperature of that object increases
If $E_{out} > E_{in}$ for an object	Internal energy of that object decreases	Temperature of that object decreases
E_{in} equals E_{out} for an	Internal energy of that object is not	Temperature of that object is not
object	changing	changing

- A cornerstone of physics is the first law of thermodynamics, which says energy is neither created nor destroyed, only altered in form.
- If an object loses energy, then some other object must gain that same amount of energy. Furthermore, because photons are just little packets of energy, the first law states that when an object emits a photon, its internal energy decreases.
- And because temperature is a measure of internal energy, the emission of a photon therefore causes the object to cool.
- Equally, if a photon hits an object and is absorbed, then the energy of the photon is transferred to the object's internal energy and the object warms.
- Radiative forcing: Imbalance between energy coming in from Sun and energy going out to space

Relationship between radiation and energy balance



All atmospheric gases have a unique pattern of energy absorption: they absorb some wavelengths of energy but are transparent to others. The absorption patterns of water vapor (blue peaks) and carbon dioxide (pink peaks) overlap in some wavelengths. Carbon dioxide is not as strong a greenhouse gas as water vapor, but it absorbs energy in wavelengths (12-15 micrometers) that water vapor does not, partially closing the "window" through which heat radiated by the surface would normally escape to space. (Illustration adapted from Robert Rohde.)

Source: NASA (n.d.)

- Greenhouse gases (CO₂, CH₄, N₂0) absorb infrared energy with wavelengths in a part of the energy spectrum that other gases, such as H₂O, do not.
- CO₂ is very strong absorber of infrared energy with wavelengths longer than 12-13 micrometers. Wavelengths of outgoing thermal infrared energy that the atmosphere's most abundant GHG —water vapour—would let escape to space are absorbed by CO₂.
- Absorption of outgoing thermal infrared by CO₂ means imbalance of E_{in} and E_{out}

3. Role of climate modelling in climate science

Surface temperature parameters



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- Surface temperature of the planet is basically determined by three parameters: the quantity of greenhouse gases in the atmosphere, the solar energy, and Earth's albedo (reflectivity)
- We can estimate Earth's temperature if we have details for these three variables
- Solar energy is largely constant (1,360 W/m²), Earth's albedo is ~0.3 (but decreasing), and greenhouse gases are quantifiable

4. Carbon cycle

Composition of atmosphere



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What are greenhouse gases?

Global net anthropogenic emissions have continued to rise across all major groups of greenhouse gases.



a. Global net anthropogenic GHG emissions 1990-2019⁽⁵⁾

Source: IPCC (2022)

Greenhouse gases and Global Warming Potential (GWP)

Global Warming Potential (GWP) concept:

- The heat absorbed by any greenhouse gas in the atmosphere, as a multiple of the heat that would be absorbed by the same mass of CO₂.
- GWP depends on gases:
 - Concentration in atmosphere
 - Strength of their absorption of infrared radiation
 - Longevity in atmosphere

Long-term fate and Global Warming Potential of different GHGs

Table 5.1 Greenhouse-gas metrics

Species	Lifetime	Global warming potential	Increase in abundance since pre-industrial times
Carbon dioxide	500 years	1	130 ppm
Methane	12.4 years	32	1.1 ppm
Nitrous oxide	109 years	260	75 ppb
Halocarbons	Years to millennia	100s to 1,000s	

Water vapour (H₂O)



 Air can hold ~7% more water vapour per 1°C warmer atmosphere





Daly (2022)



Global Atmospheric Concentrations of Methane Over Time

Data source: Compilation of five underlying datasets. See www.epa.gov/climate-indicators for specific information.

For more information, visit U.S. EPA's "Climate Change Indicators in the United States" at www.epa.gov/climate-indicators.

Source: US EPA (2022)

Agriculture and biogenic methane

Atmospheric residence times for a pulse of CO2 and CH4



Nitrous Oxide (N₂O)

Chart - Trends in atmospheric concentrations of CO2 (ppm), CH4 (ppb) and N2O (ppb), between 1800 and 2017



The global carbon cycle





Figure 5.4 Atmospheric carbon dioxide over the past 850 million years, in ppm. The dotted line is the atmospheric abundance Source: Dessler (2021) in 2020, 410 ppm.

How are humans changing carbon cycle?

ocean

Figure 5.5 Diagram of the carbon cycle as perturbed by humans. Red arrows show net flows of carbon caused by human activities. Red arrows A, B, and C represent deforestation, enhanced absorption of carbon by the land biosphere, and enhanced absorption of carbon by the ocean, respectively.

Source: Dessler (2021)

biosphere

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Atmospheric CO₂ 1960 - Present

Figure 5.7 The amount of carbon dioxide released into the atmosphere by human activities (orange line) and the observed annual increase in atmospheric carbon dioxide (blue line). Emissions data are described by Friedlingstein et al. (2019) and were downloaded from <u>www.icos-cp.eu/global-carbon-budget-2019</u>, retrieved June 9, 2020. Observations of atmospheric carbon dioxide were obtained from the NOAA Earth System Research Laboratory/Global Monitoring Division and were downloaded from <u>www.esrl.noaa.gov/gmd/ccgg/trends/data.html</u>, retrieved October 11, 2020.



5. Forcings, feedbacks and climate sensitivity

Radiative forcing – Greenhouse gases

- Increase in atmospheric CO₂ from 280 to 418 ppm between 1750 and September 2022. Corresponds to radiative forcing of +2.5 W/m²
- Because of long lifetime of CO_2 a fraction of radiative forcing the Earth is experiencing now is due to CO_2 emitted in the early 1800s.
- Increases in methane (CH₄), nitrous oxide (N₂O), halocarbons and surface ozone (O₃) (bad ozone), between 1750 and 2018 produced radiative forcings of +0.54, +0.19, +0.38 and 0.35 W/m² respectively.





Source: Dessler (2021)

Radiative forcing

Total net forcing

• Virtually all radiative forcings are linked to humans

- Radiative forcing: E_{in} > E_{out} leads to warming planet
- Current estimates that E_{in} > E_{out} by 0.8 ^a W/m². This means that planet has warmed enough in last 250 years to erase +1.7 W/m² of radiative forcing



Figure 6.3 Radiative forcing caused by human activities between 1750 and 2018. The error bars indicate the uncertainty of the estimate.

Climate feedbacks

- Feedbacks do not initiate warming but increase (positive feedback) or reduce (negative feedback) initial warming
- Fast feedbacks occur rapidly from a change in surface temperature and are important for climate change over coming century
- **Positive feedbacks:** ice albedo and water vapour feedbacks



Figure 6.8 The ice-albedo feedback loop.



Impact of fast feedbacks

- Climate modelling estimates predict a warming of ~1.2
 °C in response to doubled CO₂ BEFORE any feedbacks
- The warming AFTER feedbacks have been amplified: are 2.4 - 4.8 °C
- Doubled CO₂ (climate sensitivity) corresponds to radiative forcing of +4 W/m²
- Climate sensitivity best estimates are 0.75 °C (W/m²)
 i.e. 4 W/m² x 0.75 °C (W/m²) = 3 °C

6. Causes of climate change

Greenhouse Gases: Evidence of Warming

- Robust theory: physical scientific basis for why increased GHGs warm planet
- Geological and climatological records support theory: geological evidence of connection between GHGs and climate (e.g. Paleocene-Eocene 55-60 million years BP)



Source: NASA (n.d.)

Greenhouse Gases: Evidence of Warming

- Computer simulations of climate provide support: Climate model simulations that exclude the observed increase in GHGs don't show an increase in temperature over the second half of the 20th century
- **Fingerprints of warming:** Different warming mechanisms have unique signatures in pattern of warming



7. Future climate change predictions, including for Ireland

SSPs and future climate change

b) Contribution to global surface temperature increase from different emissions, with a dominant role of CO₂ emissions Change in global surface temperature in 2081-2100 relative to 1850-1900 (°C)



Source: IPCC (2021)

Each increment of global warming causes larger changes in regional mean temperature, precipitation and soil moisture



Source: IPCC (2021)

Source: IPCC (2021)

Total cumulative CO₂ emissions taken up by land and oceans (colours) and remaining in the atmosphere (grey) under the five illustrative scenarios from 1850 to 2100



Source: IPCC (2021)

Temperature Projections 2041-2060

- Projected temperature change for
 - a) Low-medium emission scenario (RCP4.5) and
 - b) High emission scenario (RCP8.5).
- Future period defined as 2041–2060, and is compared with the past period, 1981–2000.
- Numbers included on each plot are the minimum and maximum projected changes, displayed at their locations (<u>Nolan and</u> <u>Flanagan, 2020</u>).



Extreme Temperature Projections 2041-2060



- Projected changes in midcentury extreme temperature:
 - a) top 5% of daily maximum temperatures (warm summer days) and
 - b) bottom 5% of dailyminimum temperatures(cold winter nights).
- In each case, the future period, 2041–2060, is compared with the past period, 1981–2000.
- Numbers included on each plot are the minimum and maximum projected changes, displayed at their locations.

Seasonal Precipitation Projections 2041-2060

- Mid-century seasonal projections of average precipitation (%) for
 - a) RCP4.5 and
 - b) RCP8.5 scenarios.
- In each case, the future period, 2041– 2060, is compared with the past period, 1981–2000.
- The numbers included on each plot are the minimum and maximum projected changes, displayed at their locations





Wet and Very Wet Days Projections 2041-2060

- Projected changes (%) in mid-century number of annual
 - a) Wet days (precipitation >20mm day) and
 - b) very wet days (precipitation >30mm day).
- In each case, the future period, 2041–2060, is compared with the past period, 1981–2000.

Seminar summary

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Thank you – Questions welcome

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